GPS Precise Tracking Of Topex/Poseidon: Results and implications

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Abstract .................................................................................................................. 1

Introduction ............................................................................................................ 2

Institutional Roles .................................................................................................... 3

Experiment Goals .................................................................................................... 4

System Design .......................................................................................................... 5

The Global Positioning System ................................................................................. 5

The GPS Flight Receiver ............................................................................................ 7

The Global Tracking Network .................................................................................... 7

The GPS Operations Center ....................................................................................... 8

Solution Strategies .................................................................................................... 8

topex/poseidon Dynamic Models ............................................................................. 9

GPS Dynamic Models ............................................................................................... 10

Earth Models ............................................................................................................. 10

GIPSY-OASIS Solution Scenario, Reduced Dynamic Processing ......................... 10

Tuning Stochastic Acceleration Parameters ............................................................ 13

MSODPI, Gravity Tuning ......................................................................................... 13

SLR/DORIS Solutions ............................................................................................. 14

Orbit Quality Assessment ......................................................................................... 14

Reduced Dynamic Internal Tests .......................................................................... 15

Postfit Residuals ......................................................................................................... 15

Orbit Overlap ............................................................................................................ 15
External Tests ........................................................................................................ 16

Comparison with NASA Precise Orbit Ephemeris (POE) ......................... 16

Altimeter Crossover Analysis ............................................................................. 17

Geographical Error Distribution ...................................................................... 20

Spectrum Dynamic minus Reduced Dynamic Altitude ................................. 22

Gravity Tuning .................................................................................................... 22

Remaining questions ......................................................................................... 24

Topex/Posidon GPS Antenna............................................................................... 24

Implications for the Future ................................................................................ 26

Conclusions .......................................................................................................... 27

Acknowledgment ................................................................................................ 28

References ........................................................................................................... 28

Tables ................................................................................................................... 34

Figure Caption ..................................................................................................... 45

Figures ................................................................................................................ 48
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ABSTRACT

A "reduced dynamic" filtering strategy that exploits the unique geometric strength of the Global Positioning System to minimize the effects of force model errors has yielded orbit solutions for TOPEX/Poseidon which appear accurate to better than 3 cm (1σ) in the radial component. Reduction of force model errors also reduces the geographic correlation of the orbit error. With a traditional dynamic approach, GPS yields radial orbit accuracies of about 5 cm, comparable to the accuracy delivered by satellite laser ranging and the DORIS Doppler tracking system. A portion of the dynamic orbit error is in the JGM-2 gravity model; GPS data from TOPEX/Poseidon can readily reveal that error has been used to improve the gravity model. The operational aspects of a precise GPS tracking system have also been assessed. The TOPEX/Poseidon experience suggests that automated GPS techniques can bring the cost of precise orbit determination within that of conventional tracking systems.
INTRODUCTION

In the mid-1980s the TOPEX/Poseidon team rejected (Fuc et al., this issue) agreed to develop and fly an experimental Global Positioning System receiver to test the ability of GPS to provide precise orbit determination (POD) by an unconventional new technique (Melbourne et al., 1994). The GPS receiver aboard TOPEX/Poseidon tracks the dual 1.-band radio signals from a constellation of 24 GPS satellites, collecting navigation data from up to six satellites at once. Since the orbits and clock offsets of the GPS satellites are known (they are broadcast by the GPS satellites) the receiver can determine its position and time (four unknowns) geometrically at any instant with data from only four satellites. It is this extraordinary geometric strength that distinguishes GPS as a tracking system. Such ground-based systems as SLR (satellite laser ranging) and DORIS (Doppler Orbitography and Radio positioning integrated by Satellite) typically provide measurements in just one direction at a time and may have substantial coverage gaps; they must therefore rely on models of satellite trajectories (derived from models of the forces acting on the satellite) to recover 3-D information.

With a technique known as reduced dynamic tracking (Wu et al., 1991; Yunck et al., 1990, 1994) we can exploit the 3-D geometric strength of GPS to minimize dependence on dynamic models and, in theory, achieve a superior orbit solution through an optimal synthesis of dynamic and geometric information. A variation on that technique called kinematic tracking can yield a precise solution almost entirely by geometric means.

Conventional dynamic POD depends on precise models of the forces acting on the satellite to describe the trajectory. In a dynamic solution the estimated parameters will typically include the satellite initial state (position and velocity) and a few quantities describing the force models (e.g., drag coefficient). These are adjusted to yield a solution that best fits the observations, but that solution will necessarily have errors arising from errors in the force
models. With GPS tracking, the model errors can be observed in the 3-D residuals between the orbit solution and the observations. This residual information can then be applied in a point-by-point geometric adjustment of the satellite position to give the reduced dynamic solution (Fig. 1). Differences between dynamic and reduced dynamic solutions can expose the model errors and allow us to study their geographical and spectral distribution. Alternatively, parameters describing the gravity field can be adjusted in a dynamic GPS solution to improve (or tune) the gravity model with an unprecedented degree of global strength.

Institutional Roles

This work involved a collaboration between groups at the Jet Propulsion Laboratory (JPL), the Center for Space Research (CSR) of the University of Texas at Austin, and a scientist visiting JPL from the Institut Geographique National (IGN) in Paris. The JPL team focused on refining the reduced dynamic strategy, while CSR, which has long experience in dynamic estimation with S1.R, adapted their software for dynamic POD and gravity tuning with GPS data (Rim, 1992). Although the JPL and CSR analysis systems were developed independently (Wu et al., 1990; Webb et al., 1993; Rim, 1992), they share some common models. Comparisons between orbits produced with each system serve as an important validation test.

IGN has expertise in the DORIS system and worked closely with JPL to adapt JPL’s analysis software to process DORIS data. CSR modified its software independently to process DORIS data as well (Watkins et al., 1992). In addition, complementary efforts are underway at the Goddard Space Flight Center (GSFC) and CNES (Centre National d’Etudes Spatiales) in Toulouse to produce the official precise orbits with S1.R and DORIS data (Tapley et al., this issue).
Experiment Goals

The major goals of the GPS experiment are: (1) evaluate the accuracy and operational potential of GPS for tracking Earth satellites; (2) provide data that includes the GPS-based orbit solutions, calibration data, and reference frame ticks for post-experiment use by the Project (Born et al., this issue; Christensen et al., this issue); and (3) provide production PSPPOD technology for possible conversion to an operational system.

In the 1980s, covariance analysis suggested that an accuracy of 5 to 10 cm might be achieved if data from 6 globally distributed GPS ground receivers were used together with the flight data to solve for the TOPEX/POSEIDON orbit (Wu and Ondrasik, 1982; Yunck and Wu, 1986; Wu et al., 1987). We therefore adopted “better than 10 cm RMS in altitude” as a formal goal for the experiment. Analysis further showed that ionospheric calibration with dual frequency GPS data would also be needed. We should note that the flight receiver developed for the experiment cannot receive both frequencies when the GPS security feature known as anti-spoofing (AS) is active. It was therefore necessary to arrange, with the Department of Defense (DoD), to have AS off for nine 1()-clay periods in the first year of the mission to ensure an adequate data set for analysis. Future receiver designs could avoid this problem by adopting, either a GPS decryption capability or advanced codeless tracking techniques.

Our objective in evaluating the operational potential the GPSPOD system is to see if GPS can be a cost-effective alternative to existing precise tracking systems. Measures of operational performance include time delay in producing and validating the precise orbit products, reliability of the system, and cost of operation.

Recognizing the potential to support the TOPEX/POSEIDON Project more formally, we also set out to collect, edit, and archive all data over the experiment lifetime (1 year for the
flight receiver & 2 years for the ground network); 2) tune the gravity model to improve, the ocean geoid at wavelengths >1000 km; and 3) make available to the Project the most precise orbits for use in oceanographic studies and for altimetric calibration at the verification sites.

SYSTEM DESIGN

The GPS tracking system consists of four segments: the GPS constellation, the flight receiver, a global network of GPS ground receivers, and a central monitor, control and processing facility (Fig. 2). The POI strategy requires continuous tracking of the visible GPS satellites by ground and flight receivers. Data from all receivers are brought together and processed in a grand solution in which the TOPEX/Poseidon mbits, all GPS orbits, receiver and transmitter clock offsets, carrier phase biases, and a number of other parameters are estimated. Simultaneous sampling at all receivers (which may be achieved by later interpolation) eliminates common errors, such as clock dithering, which is a feature of another GPS security feature, known as selective availability (SA). In the end, TOPEX/Poseidon position and velocity are determined in a reference frame established by key sites in the global network; those sites are known absolutely with respect to the geocenter to about 2 cm in the International Terrestrial Reference Frame.

The Global Positioning System

Figure 3 depicts the GPS constellation, which is controlled from Falcon A1-B near Colorado Springs. The constellation consists of 24 GPS satellites in 12-hoLl (20,200-km altitude) circular orbits (Milliken and Zoller, 1978; Spilker, 1978). The satellites are distributed in six orbit planes inclined at 55°, with a nodal separation of 60°. Each satellite broadcasts navigation signals on two L-band frequencies: 1.57542 GHz (1.1) and 1.2276 GHz (1.2). The corresponding carrier wavelengths are approximately 19 and 24 cm. The two
Carrier phase = range + clock offset + troposphere + noise
Pseudorange = range + clock offset + troposphere + noise

The observations are given in a simplified form by
between the transmitter and receiver and from various instrumented phases. The troposphere
process and, hence some arbitrary bias resulting from the unknown number of whole cycles
quantity as pseudorange, with two distinct differences: it is about one hundred times more
delay has been removed by dual-frequency combination, carrier phase measures the same
can be seen in an additional delay caused by the earth's atmosphere. After the troposphere
measurement, however, is controlled by various other factors. The ground receivers, for
measurement purposes, thus the offset between the transmitter and receiver clocks. The pseudorange
accuracies, the range between the phase centers of the GPS satellites and receiver

Pseudorange is the part of the code range observations, and the primary GPS data type.

receiver as part of the code range observations, and the primary GPS data type.

The process of squaring and band center phase measurements, which are recovered by the
known as GPS time, selecting the best returns, and other information of value to the user. For
phases that are necessary, which provides accurate GPS orbits, clock offsets from a fixed standard
correction/erasure of C/A-code, finally, both L-band signals are further modulated by a 50
quadrature (90°) one of phase from the P-code) by a less precise ranging code known as the
combination separation of the received GPS signals. The L1 signal is also modulated in
measurement of the pseudorange. Each signal broadcasts a unique code orthogonal to the others,
where the transmitted phase (according to the transmitter clock) of each bit is known, this gives a
P-code. The receiver measures precisely and unambiguously the arrival time of each code bit
[2]. L-band carrier is modulated with a precise pseudo-random ranging code known as the
receptions are clearly visible within a vertically centered hexagonal pattern of field of view.
any point on the earth's surface, or in space below 3000 km, typically 5 to 9 GPS
exceeds roughly 3000 km beyond the limb of the earth as viewed from the GPS satellites. All
accuracies are used to calibrate the tropospheric delay. The beamwidths of the GPS signals

6
A more detailed description is given by Wu et al. (1990).

**The GPS Flight Receiver.**

Figure 4 is a sketch of the TOPEX/Poseidon spacecraft, showing the locations of some subsystems and flight instruments. The GPS antenna is atop a 4.3-m mast, above the main body of the satellite, to suppress reflected signals from the TDRS high-gain antenna and other prominent surfaces. The GPS Demonstration Receiver (GPSDR), an early version of the Motorola Monarch™ (not visible) tracks up to six GPS satellites concurrently, measuring the phase of each carrier at 1-sec intervals and pseudorange at 10-sec intervals. Measurement noise on the ionosphere-free observables, including instrumental thermal noise and multipath effects, is about 5 mm for phase and 70 cm for pseudorange. For details on the flight receiver see Zieger et al. (1994).

**The Global Tracking Network.**

Figure 5 shows the primary ground sites used in the experiment. These will be part of the International GPS Service (IGS) set to begin providing high accuracy GPS data products to scientific users in 1994, under the auspices of the International Association of Geodesy (Neill et al., 1993). For TOPEX/Poseidon fewer than a dozen sites are needed to obtain full accuracy because of the ample common GPS visibility between the satellite and the ground sites. For GPS ground programs (which now achieve a weekly geocentric station location precision of about 1 cm), 20-40 sites are sometimes required (Blewitt et al., 1993).
The GPS Operations Center

All transactions involving GPS data and POD products flow through the operations center, which automatically retrieves data from all GPS sources—about 8 Mbyte/day from the flight receiver and 1 Mbyte/day from each ground site. The center monitors and controls the ground and flight receivers and initiates actions to repair system faults. The Rogue™ and Turbo-Rogue™ ground receivers can store their data for, in most cases, up to 12 days to protect against communication outages. In the first 6 months of experimental operations we acquired 99% of the possible data from the flight receiver when GPS anti-spoofing was off, and about 95% from the ground receivers.

Precise GPS-based mbits for TOPEX/Poseidon are now produced at JPL with 30-hr data arcson 2.4-h' centers, providing 6-hr Over1aps for comparisons. Those orbits and statistical quality measures are available about 8 hrs after all data for a 30-hr arc arc received. External release of the orbits occurs about 3 days after the end of each 10-day mbit repeat cycle. Processing of the orbits is automated and data driven. Once the analysis process is initiated on the workstation it runs continuously, around the clock, with no operator attention except to deal with rare anomalies. The process wakes up every 3 hrs to see if the data for a given arc have arrived. When the required data are there, processing for a 30-hr arc begins.

SOLUTIONS STRATEGIES

Here we compare TOPEX/Poseidon precise orbits computed by three groups: JPL, CSR, and GSFC. Each group used different analysis software, applied to one or more of three precise tracking data types: GS, DORIS and S1JR. GSFC employs a combination of S1R and DORIS data to deliver operationally the precise orbits placed on the official Geophysical Data Records (GDRs) distributed to scientists (Tapley et al., this issue). JPL and CSR have
performed experimental analysis of the GPS data, and have analyzed some combination of SLR and DORIS data as well. While their orbit estimation techniques differ in important ways, the three groups share common models for TOPEX/POSEIDON dynamics and for the positions of observing (or transmitting) points on the Earth relative to inertial space, in which the orbit is propagated. JPL's strategy is unique among the three in its use of Kalman filtering and stochastic models to permit reduced dynamic orbit determination.

**TOPEX/POSEIDON Dynamic Models**

While the analysis systems share common dynamic models, those models are realized through implementations which give slight differences in the computed ocean tides and Earth albedo (Tapley et al., this issue). All solutions, unless otherwise noted, use the JGM-2 gravity field tuned with TOPEX/POSEIDON SLR and DORIS data (Lerch et al., 1993; Nerem et al., this issue). A custom model for the solar and thermal radiation forces on TOPEX/POSEIDON was developed for the SLR/DORIS effort (Marshall et al., 1992). The thermal radiation portion of the model was not used in the JPL GPS solutions, however. These small, slowly varying dynamic model differences can be largely accommodated through the adjustment of an empirical acceleration parameter, $\vec{\alpha}$, of the form

$$\vec{\alpha} = \vec{C} + \sum_{i=1}^{2} \vec{A}_i \cos \omega_i t + \vec{B}_i \sin \omega_i t$$

where $\vec{C}$, $\vec{A}_i$, and $\vec{B}_i$ are constant vectors in the spacecraft coordinate system oriented in the nominal along-track and cross-track directions (Kaplan, 1976). The frequencies $\omega_i$ are once- and twice-per-revolution of TOPEX/POSEIDON and $t$ is time past an epoch. Solutions produced by CSR (with UTOPIA and MSfdc) and GSFC (with Geodyn) adjusted constant and once-per-revolution along-track and cross-track amplitudes, while JPL's preliminary dynamic solutions (with GIPSY/OASIS 11) adjusted twice-per-revolution terms in those components as well.
**GPS Dynamic Model**

The dynamic model for the GPS satellites contains only two components: the JGM-2 gravity field up to degree and order 12, and custom solar radiation force models known as T10 and T20 (Fiege et al., 1992).

**I<i>lll Models**

All three analysis systems use the IERS Standards set forth in IERS Tech Note 13 (McCarthy, 1993) (J'alle et al., this issue) for earth orientation and the deformation of the earth due to solid and polar tides. JPL's GPS solutions estimated polar motion and UT1 with nominal values taken from IERS Bulletin B finals or predicts, depending on the time of processing. The CSR and GSF C solutions employed polar motion and UT1 rate values determined by SLR data from Lageos (Tapeley et al., this issue).

**GIPSY-OASIS II Solution Scenario, Reduced Dynamic Processing**

JPL's computed dynamic and reduced dynamic solutions with the GIPSY-OASIS 11 analysis software (Webb et al., 1993; Wu et al., 1990). Its main components are a GPS data editor, orbit integrator, measurement model generator, and filter/smoothcr. The data editor operates on a combined set of dual frequency GPS phase and pseudorange measurements. It automatically detects outliers and carrier phase discontinuities (Blewitt, 1990). An automated executive ties the modules together, producing daily orbit solutions unattended. The system typically produces a reduced dynamic solution within 2 days of onboard data acquisition, using less than 6 CPU hours on an II'735 workstation.
The orbit integrator numerically integrates the satellite trajectory from a nominal initial state using precise models of the forces acting on the satellite. It also computes partial derivatives of the current state of the spacecraft with respect to the dynamical and epoch state parameters. The trajectory and partials are then passed to the measurement model program.

After editing, the data are compressed to 5-rein normal points and the ionosphere-free combinations of phase and pseudorange are formed. In the compression step the pseudorange data are smoothed against the carrier over the entire S-rein interval, while the phase is simply sampled at the appropriate times. Because the TOPEX/Poseidon onboard clock drifts freely with respect to the ground receiver clocks (which are kept close to UTC), we require a small interpolation of onboard phase to the appropriate sample time to ensure common mode cancellation of SA dithering. This is accomplished with a cubic fit to four 1-sec points about the desired time (WU et al., 1992). The nominal trajectory is then used to compute model GPS observables and their partial derivatives with respect to all adjusted parameters. The measurement model program then retrieves the satellite positions and partials passed by the integrator, computes the model observables, and, in addition, partial derivatives of the observables with respect to ground station position, zenith tropospheric delay, earth orientation, GPS clocks, and receiver clocks. The observable model includes relativistic effects, the Earth models discussed above, phase windup due to antenna rotation (Wu et al., 1993), antenna phase-mnter variation as a function of azimuth and elevation (Zieger et al., 1994).

Next, the filter/smoothler takes over to carry out the grand solution for the TOPEX/Poseidon and GPS states, ground site positions (five are held fixed for reference), clocks, atmospheric delays, and so on. In its simplest mode, the filter/smoothler produces the equivalent of a conventional batch least-squares solution; but to obtain a more accurate orbit, some parameters are treated as stochastic processes and adjusted at each time step in a time-sequential Square
Root Information Filter (SRI) formulation (Bierman, 1977). The parameters adjusted in our standard solution strategy are summarized in Table 1.

In these solutions, all clocks are solved for freely and independently at each S-reach time step (i.e., modeled as white noise processes with no a priori constraint, except for one at a ground station which is held fixed as a reference clock). The zenith atmospheric delay at each ground site is also adjusted at each step, modeled as a random walk which in 1 hr adds 1 cm uncertainty in the zenith delay. For the 30-hour dataacrs, the parameters of the T10 and T20 solar pressure model (Fliegel et al., 1992) are treated as loosely-constrained constants plus a small colored noise process with a 4-hour correlation time and sigma of 10% at 1-hr batch times. The estimation of the GPS orbits is essentially dynamic.

The reduced dynamic solution is produced only in the final estimation step. First, the “1011’X’Poseidon epoch state and the empirical constant and once- and twice-per-revolution accelerations (Table 3) are adjusted to convergence in a dynamic solution, which takes two passes through the filter. This dynamic solution is typically accurate to better than 20 cm (31), well within the linear regime for the final reduced dynamic adjustment. In the reduced dynamic step, adjustments are made to the X’Poseidon state and to all previously adjusted parameters except two types: the empirical once- and twice-per-revolution terms, which are now held fixed, and the constant accelerations (C in Table 3), which now become stochastic and are re-estimated at each time step to provide the local geometric corrections. The latter are modeled as first order Gauss-Markov (colored noise) processes and given a correlation time of 15 min with steady-state sigmas of 10, 20, and 20 nanometers/sec in the radial, cross- and along-track directions. It is the geometric strength of the GPS observations that allows these final stochastic adjustments to be made with high accuracy.
Tuning Stochastic Acceleration Parameters

The steady-state sigmas for the stochastic acceleration parameters were chosen through an empirical process in which solutions were generated with a range of sigmas, and the final values selected were those that minimized the RMS differences on the 6-hr orbit overlaps for several test arcs. Once chosen, they were held fixed in all processing. A better criterion might be altimeter crossover statistics, but those were not available in our earliest processing and were later reserved as an independent test of orbit accuracy (see tests below).

MSODP1, Gravity Tuning

The Center for Space Research/University of Texas at Austin used MSODP1 (Multisatellite orbit cte.termination Program) for the GPS/Topex data processing. The program has been compared against UTOPIA, the single-satellite mbt determination program used for processing S1, S1, and DORIS data, and the two have agreed at the centimeter level.

h4S01J1 uses doubly differenced phase measurements between the flight receiver and the ground stations at 30-sec intervals. For the experiments presented in this paper, no double differences between pairs of ground stations were used. All double differences were corrected for the ionosphere, and pseudorange measurements were used to compute each receiver clock offset from GPS time.

The MSODP1 uses a batch-least squares estimator implemented with a square-root-free Givens algorithm for improved numerical stability. No a priori constraints are assigned to any estimated parameter. A simultaneous solution is performed for the TOPEX/Poseidon and all GPS satellite states, along with once/revolution parameters for TOPEX/Poseidon and radiation pressure parameters for the GPS satellites. A constant zenith tropospheric delay is estimated.
at each site every 2.5 hrs, and a phase bias parameter is estimated for each combination of TOPEX/Poseidon, GPS satellite and ground receiver. One-day solution arcs were used in all cases except for the tuning of the gravity field, where 3.3-day arcs were used.

**SLR/DORIS Solutions**

The precise orbit ephemeris (POE) produced for the altimeter geophysical data records and released to the science community is computed dynamically by GSFC with SLR/DORIS data over 10 day arcs. They are released only after an extensive validation procedure (Tapley et al., this issue). We will make comparisons to these official orbits as a test of the GPS reduced dynamic orbits.

**ORBIT QUALITY ASSESSMENT**

First we describe internal consistency tests within the GIPSY-OASIS II processing system, and then compare the GPS-reduced dynamic orbit S1 with the GSFC POE solutions. Next we present altimetry crossover differences, which provide a test that is independent of all orbit determination techniques and software. Next we examine the difference between the dynamic and reduced dynamic orbits produced with GIPSY-OASIS II to obtain information on the geographically correlated orbit error and its spectral content. The CSR group has recently tuned the JGM-2 gravity model with GPS data; in the final test, dynamic solutions produced by CSR with the tuned field are compared to reduced dynamic solutions made with JGM-2.
of the overlap is shown in Fig. 7. The RMS difference is 0.88 cm in altitude, 5.70 cm cross.

for the stochastic acceleration, a sample of the orbital difference during the central 4.5 hours
are computed in the RMS comparisons. This corresponds to three times the time
read with reduced dynamic solutions, 45-min segments from each of the two solutions
in the absence of data on the other side to constrain the stochastic estimate.)

To avoid the estimation effects, increased effort on the ends of the solution was

Left: overlap

Only 0.01% of data are detected as anomalous and removed from the filtered solution.

The RMS data in general of high quality; the residuals are typically removed, phase residuals are typically
removed. Phase residuals for the high receiver are typically about 1.5

As far as anomalous quality control, the software examines positional phase and

Positive Residuals

Dynamic Initial Tests
track and 3.44 cm along track. Fig. 8 shows the average RMS overlap agreement in altitude for all overlaps for twelve 10-day cycles. The agreement is consistently below 2 cm, with an average of about 1 cm. The anomalous value for cycle 21 appears to have been caused by data outages at Goldstone while Goldstone was used as the reference clock. We have since modified the automated analysis to prevent the use of a reference clock at a station with sizable outages. Cycle 19, which produced the best agreement, was the only cycle in which no GPS satellites passed through the earth's shadow. During such eclipses the GPS force and measure.ment model errors increase noticeably. The TOPEX/POSEIDON dynamic overlap agreement (not shown) is consistently worse, giving single RMS altitude overlap differences as high as 5.6 cm and an average RMS difference of about 2 cm.

External Tests

Comparison with NASA Precise Orbit Ephemeris (POE)

Figures 9 and 10 show the RMS differences between JPL's GPS solutions (both dynamic and reduced dynamic) and the NASA POE over six 10-day repeat cycles. The average RMS radial difference was 2.68 cm for the dynamic comparison and 3.33 cm for the reduced dynamic comparison. The maximum differences in radial position at any point over all six cycles were 12.2 cm (dynamic) and 11.5 cm (reduced dynamic). We shall argue that the better RMS agreement between the two dynamic orbits is the result of common crews in JGIM-2 and the non-gravitational force models, which are partially removed in the reduced dynamic solution.

In comparing the JPL dynamic and reduced dynamic orbits against the NASA POE, a bias in the mean of the z coordinates of the Greenwich Reference Frame (pseudoearth-fixed) of about 3 cm was noticed. This bias varies slightly from cycle to cycle and day to day.
Most of the mean differences in the x and y coordinates can be attributed to errors in JGM-2, as suggested by the much smaller differences in the dynamic solutions (Table 2) and the offset predicted by the JGM-2 covariance shown in Plate 1b (see discussion below on the geographical correlation of the radial errors). The mean z bias remains essentially unchanged whether a dynamic or reduced dynamic mbit is used in the PO1 comparison. The z bias also appears in comparisons of the J'II' mbits to CSR orbits computed with either GPS or S1. R/DORIS data (Table 8). We note that recent determinations of the geocenter from GPS ground data only have obtained decimeter level accuracy in the z component (Vigue et al., 1992). Inclusion of TOPEX/Poseidon data in geocenter solutions has improved the observability of this component to about the centimeter level (Tapley et al., 1993b; Malla et al., 1993). Although the observed z bias between the J'II' and other orbits does not appear to reflect a limitation of GPS tracking, we have yet to identify its source and continue to look for it. A 3 cm translation in z reduces the RMS differences by about 3 mm.

If we assume that the errors in the reduced dynamic orbits and the PO1 mbit are uncorrelated we can attempt to allocate the 3.33 cm RMS difference. An equal allocation would yield an RMS radial error of 2.35 cm for both solutions. However, using altimeter crossover analysis and the geographical distribution of errors, we will argue that the errors between the two mbits are largely uncorrelated and that the reduced dynamic orbit error is somewhat smaller.

**Altimeter Crossover Analysis**

A key method for assessing the relative radial accuracy of different orbits relies on altimeter data collected by the spacecraft TOPEX/Poseidon. It carries two nadir-pointing radar altimeters that measure the range to the seafloor with an uncertainty of less than 4 cm RMS (Vigne et al., Ibis issue). These range measurements can be used together with the precise radial orbit solution to determine the geocentric height of the seafloor. At the points in the ocean
where the satellite ground tracks intersect ascending and descending passes, two such determinations of sea height can be made. In the absence of errors in the radial component of the orbit and in the media corrections to the altimeter range, the height difference at the crossing point location is a measure of the true variability of the ocean surface.

Crossover observations from eight separate 10-day repeat cycles of the TOPEX/Poseidon ground track were used for this analysis (Aviso, 1993). Since there is a range bias of about 20 cm between the two altimeter systems (Christensen et al., this issue), we used only the data from the U.S. dual-frequency altimeter. All standard environmental and sc:ix-state corrections were applied and editing was performed based on the data flags provided with the crossover geophysical records. As crossovers may occur days apart, corrections for ocean dynamic effects, such as those attributable to tides (Cartwright and Ray, 1990) and atmospheric pressure loading, were also applied. A confounding factor is the unmodeled sea height variation from changes in ocean currents and errors in tide models and media corrections. To mitigate the effects of current variations, we restricted our analysis to crossovers occurring within the individual cycles. Table 5 lists the global crossover statistics for the GPS-reduced dynamic orbits and for the two precise orbits provided with the merged GDR products. Over 35,000 individual crossovers occurring in the period from January 30 to May 19, 1993 are represented in the global statistic.

The actual radial orbit error is difficult to quantify based on these statistics since the residuals also contain errors in the media corrections and unmodeled oceanographic effects. A large portion of the tidal and atmospheric pressure signal has been removed with global models, but a sizable signal remains. In order to address this difficulty, we have segregated a small number of crossovers from the original global data set using a highly restrictive set of geophysical editing criteria (Table 6). (No outlier editing was performed since it is impossible to guarantee they do not result from large excursions in the orbit error.) These
editing criteria are designed to reduce the ocean variation component of the crossover residuals while maintaining a global distribution of data. To the extent that the geophysical and environmental corrections being interrogated are not correlated with the orbit error, this approach should help to better isolate the mbn error contribution.

Table 7 lists the global crossover statistics for the data remaining after the restrictive editing. Note that while only 3% of the original data remain, there are still over 1000 globally distributed observations (Fig. 11). The variance (energy) has been reduced by over 50%, corroborating that the scatter of the original data set primarily reflects contributions from non-orbit sources. Assuming that the residual variabilities are uncorrelated in a global sense on ascending and descending tracks, one could infer that the radial orbit error is less than 5 cm RMS ($7.03/\sqrt{2}$), regardless of the orbit solution under consideration. Contained in this figure is some residual error from the geophysical corrections and instrumental effects, as well as orbit error. On the other hand, if there are large stationary orbit errors that are highly correlated cm ascending and descending passes, an extreme example is an error in the overall scale of the orbit—then the crossover observations cannot observe them. Despite these caveats, the crossover statistics provide a powerful and independent tool for measuring orbit consistency and for gauging improvement. In this context, we note that the GPS-based reduced dynamic orbits yield the lowest crossover residuals. in particular, the variances of the crossover populations in both tables (cf. Table 5, Table 7) are about 10 cm$^2$ lower with the reduced dynamic orbits, suggesting a consistent reduction in TOPEX/Poseidon orbit error. If we assume that 3-4 cm of error remains from residual errors in the environmental and geophysical corrections and from ocean variability (a purely speculative number), then we can estimate that the GPS reduced dynamic orbit has a radial RMS error of 2-3 cm while the various dynamic orbits have radial RMS errors of 3-4 cm.
Past ocean altimetry missions have been plagued by what are known as geographically correlated orbit errors—that is, orbit solutions that are consistently biased in different geographic regions (T'apley and Rosborough, 1985). Such errors can confound the interpretation of altimetry data by mimicking large-scale features in the ocean topography from which circulation estimates are derived. Geographically correlated orbit errors are most commonly associated with errors in the gravity model, although coordinate system offsets and other factors may also play a role. A pre-launch covariance study by Rosborough and Mitchell (1990) showed that kinematic and reduced dynamic orbits, by reducing dependence on force model, in general, could virtually eliminate the geographic correlation in the gravity-induced TOPEX/Poseidon orbit error at large-scales. We have corroborated this result using the actual (iL'S-based) orbits for TOPEX/Poseidon.

The differences between G1'S-based dynamic and reduced dynamic TOPEX/Poseidon orbits over three 10-day periods beginning March 10, March 20, and April 1, 1993, respectively, have been analyzed in terms of the geographical distribution of errors (Christensen et al., 1993). This analysis suggests that the pre-launch gravity model, JGM-1 (Nerem et al., 1993), introduces geographically correlated errors having a strong meridional dependence. These errors can be approximated by a large-scale positive anomaly in the Indian Ocean and a large-scale negative anomaly in the eastern Pacific Ocean (Plate 1a). The global distribution and magnitude of these geographically correlated errors are consistent with pre-launch covariance analysis; moreover, the estimated and predicted global RMS error statistics are also in close agreement at 2.3 and 2.4 cm rms, respectively (Christensen et al., 1993).

The most compelling evidence that this anomaly is attributable to an error in the JGM-1 gravity model can be seen in Plate 1b, depicting the global distribution of the mean orbits.
height differences between two GPS dynamic orbits produced with the JGM-1 and JGM-2 gravity models. JGM-2 is basically the JGM-1 model tuned with TOPEX/Poseidon S1 DORIS data gathered from September 20, 1992 through February 18, 1993 (Nerem et al., 1993). Note that these differences come entirely from the gravity model since this is the only difference between the two cases. This figure is remarkably similar to Plate la, with the exception that the meridional variation is smaller and there is less track iness, i.e. less disparity among neighboring tracks. It is important to note that no GPS data were used to obtain JGM-2, so it is impossible that a geographically correlated error in the GPS tracking system would appear as an alias in JGM-2.

Repeating the GPS analysis with the JGM-2 gravity model suggests that a portion of the meridional dependence observed in JGM-1 still remains (see Plate 1 c). Though JGM-2 is a clear improvement over JGM-1, a measurable amount, 1.2 cm rms, in the differences between reduced-dynamic and dynamic orbits determined with JGM-2 persists. The salient features in this figure have also been identified in comparisons between the GPS reduced dynamic orbit and the NASA POI (also based on JGM-2), though the interpretation of this result in the context of gravity error is complicated by evident coordinate system differences. (The apparent shift between the NASA POI and the GPS-based orbits along the spin axis is discussed elsewhere in this paper.) Further comparisons between various dynamic and reduced-dynamic orbits should help to separate and identify the sources of the geographically correlated errors. Note that, as illustrated in Plate 1 b, classical dynamic orbit determination is also capable of observing small modeling errors, such as those introduced by the pre-launch JGM-1 gravity model. To accomplish this, however, the force models must be tuned with comprehensive tracking data from many orbits.

It has long been recognized that differential GPS data can be used with dynamic orbit determination techniques to improve the earth's gravity model (Bertiger et al., 1992; Tapley
et al., 1993). With an improved gravity model, G1'S-based dynamic orbits will improve and, for TOPEX/Poseidon, should approach the accuracy of reduced-dynamic orbits. (Properly weighted, however, the reduced dynamic orbits will in theory remain superior, if only by a small amount, by reducing non-gravitational and residual gravity model errors.) For orbiters at much lower altitudes, gravity and aerodynamic forces are extremely difficult to model, and the reduced dynamic technique will be crucial if sub-decimeter accuracy is needed.

Spectrum Dynamic minus Reduced Dynamic Altitude

Figure 12 gives the amplitude spectrum of the dynamic-minus-reduced dynamic altitude over 10 days. The spectrum is typical of gravity model error in a dynamic solution, which, because of the daily rotation of the field, generates a suite of tones at 1/rev ±m/day (Rosborough, 1986). The ±m/day tones in the spectrum may also include artifacts from the daily orbit fits spliced together to form a 10-day solution. Notice that nearly all of the energy is at frequencies below twice/rev. Since we know of no forces that can cause significant high frequency satellite motion, the reduced dynamic process noise constraints (the time constant and steady-state sigmas) have been set to suppress high frequency corrections. Experiments have shown that when those constraints are relaxed and we go to a more fully kinematic solution, the high frequency components increase, but only slightly, and almost entirely as a result of tracking measurement noise rather than real satellite motion.

Gravity Tuning

A preliminary tuning of the JGM-1 prelaunch gravity model has been performed at CSR by augmenting the JGM-1 gravity coefficients and their associated covariance with new information equations from twenty 10-day repeat cycles of DORIS and SLR data, and four cycles of GPS data, from TOPEX/Poseidon. This field, referred to as T1G-3, is discussed in
detail in Tapley et al. (1993). The SLR and DORIS tracking data were processed in 10-day arcs, while, for computational efficiency and to reduce GPS satellite dynamic error, the GPS data were processed in 3.3-day arcs. For both data sets, daily once/revolution empirical accelerations in the along-track and cross-track components, along with daily mean along-track accelerations, were adjusted.

The RMS of the mean, or constant component of the geographically correlated gravity errors for TOPEX/Poseidon, predicted from the covariance, has been reduced from 1.6 cm for the JGM-2 model to less than 1 cm. The peak has been reduced from 2 cm ill JGM-2 to less than 1 cm in JIG-3. We attribute this to the additional geographic coverage provided by G1'S tracking, since an alternate field (JIG-3A), obtained with S1 , Rand DORIS data only, yielded correlated errors approximately 50% larger. Separate gravity tuning solutions made with four cycles of GPS data only were generally comparable to or better than those made with 20 cycles of S1 , Rand DORIS data.

The dynamic orbits for TOPEX/Poseidon computed with JIG-3 using S1 , Rand DORIS data give better agreement with the J' ] reduced dynamic orbits (Table 8) than the corresponding orbit using JGM-2. The radial RMS agreement is currently 2.5 cm over the nine cycles examined, after removing an azimuth bias of 1.5 - 3.0 cm, with little spatially correlated signal. The three-dimensional RMS differences have also been improved for most cycles (Table 8). It should be recalled that non-gravitational orbit errors are also included in these trajectory differences, although such errors are smallest in the radial component.

The geographically correlated portions of the differences between dynamic orbits computed with JGM-2 and JIG-3 are similar to those between JGM-2 dynamic and the J' ], reduced dynamic orbits, indicating that the reduced dynamic solution and the tuned gravity field solution have converged to a similar result, although they were arrived at through quite
different filtering strategies. We regard this as evidence of both an improvement in the gravity model and the efficacy of reduced dynamic filtering.

REMAINING QUESTIONS

*Topex/Poseidon GPS Antenna*

Early in the GPS data analysis, a bias of \(-6\) cm in the radial direction was noted between reduced dynamic and dynamic orbits. Reduced dynamic tracking is sensitive to the position of the phase center of the GPS receiving antenna, and the center of gravity (CG) of the satellite is inferred from attitude information and measurements of the antenna position made before launch. The dynamic solution, on the other hand, is highly sensitive to the satellite CG since that is the reference point for the orbital dynamics. By adding a radial bias parameter to the dynamic state estimate and fixing the result in the reduced dynamic estimate the bias between the reduced dynamic and dynamic orbits is eliminated. The formal error on the radial bias estimate is about 3.3 mm with 30 hrs of data. The mean of the radial bias estimate, determined with 30-hr arcs over more than 5 months from cycle 17 to cycle 33, is 59.9 mm, and its standard deviation is 4.8 mm (Fig. 13). The maximum value occurs during a yaw ramp in the *TOPEX/Poseidon* attitude control, which might suggest a mismodeling of the yaw ramp. The minimum occurs during a planned *TOPEX/Poseidon* orbit maintenance maneuver. Figure 14 shows the time series during a period in which none of the GPS satellites are in eclipse. There are known errors in the force model for the GPS spacecraft which are larger during eclipse. There are also known errors in the measurement model for GPS due to attitude mismodeling during eclipse. Both of these models can be improved. Note the smaller standard deviation of 1.6 mm during this period.
"There are two known systematic errors in the phase observable which when corrected change the antenna offset by 1.05 cm (leaving a total offset of about 5 cm) and affect the orbits by an RMS of 1-3 mm. Together, they can be as large as 1.3 cm. The first is in the receiver decimation filler, which results from the signal’s Doppler offset, and the second is in the Costas loop, which results from the Doppler rate. These phase errors can be corrected with the following simple formula:

$$\Delta \rho = \frac{1}{2} T_c \ddot{\rho} + \frac{1}{\omega_n^2} \dot{\rho}$$

where $\dot{\rho}$ and $\ddot{\rho}$ are the range rate and range acceleration corresponding to the Doppler rate, $T_c = 0.976 \times 10^6$ sec is the code correlation period; $\omega_n$, the natural frequency of the Costas loop, is 29.3 radians/sec for the 17 Hz loop noise bandwidth of the receiver.

The mechanical position of the antenna was measured in satellite coordinate to < 1 mm before launch. It agrees to the nominal location as specified on the drawings within 2 mm in the z component and within 8 mm 3D. An anomaly exists somewhere in the overall mode 1 of the GPS observable. Although an error in the satellite measurements would explain the results, we have all but ruled that out based on other evidence and as yet have no satisfactory explanation for the apparent antenna bias. ‘There is a slim possibility that the offset could result from incomplete knowledge of the phase center of the GPS transmitters. A preliminary analysis, however, shows extremely small variation of the GPS satellite phase center with look angle. A Block-2 GPS transmit antenna from the qualification model satellite has been obtained from the USAF Space Command and is being calibrated at JPL. Meanwhile, we continue to estimate a phase center offset, even though it is now well characterized.
IMPLICATIONS FOR THE FUTURE

Results from TOPIX/Poseidon show that pre-launch covariance studies were quite accurate and thus lend confidence to predictions made by similar studies for future missions. Figure 15 gives results from such a study performed several years before launch (Wu et al., 1991). The assumptions (Table 9) were in many ways inconsistent with what has been done in the actual 30-hr J-array solutions. The available computing power at the time of the study was meager by today’s standards, requiring the use of much shorter data arcs. To compensate, we assumed a 2 m a priori error on the GPS orbits and a low pseudorange noise of 5 cm. The estimated 2-3 cm RMS radial error for the reduced dynamic solution with JGM-2 is plotted (point $x$) in Fig. 15. In addition, the typical RMS difference with SLR/DORIS solutions (with JGM-2) is plotted. Somewhat fortuitously, the artificial compensation has proved reasonably accurate, and the agreement with the estimated actual error is within a centimeter.

An example of a possible future mission is taken from the Earth Observing System (EOS), a suite of scientific Earth probes planned to fly at about 700 km beginning in the late 1990’s. Because dynamic mode errors can grow large at that altitude, a purely kinematic analysis is presented. This time the reference silt error is reduced to 3 cm per component and the number of flight receiver channels is increased to track all satellites in a hemisphere, which increases the geometric strength compared to TOPIX/Poseidon. Other assumptions that differ from the TOPIX/Poseidon covariance analysis are given in Table 10. Figure 16 shows the predicted altitude error as a function of data arc length for several different GPS data combinations. The data type called “carrier-quality range” is a fictitious pseudorange measurement having the precision of carrier phase, and serves to establish a performance bound. With data arcs longer than 20 hours, all scenarios yield about 2.5 cm RMS radial errors for kinematic tracking, which is completely independent of dynamic model error.
To see what might be done with even greater geometric strength we present a study of the Space Shuttle at 300 km in which we open up the flight receiver field of view to the full sky (each Shuttle is equipped with GPS antennas top and bottom to permit this). Typically, there will be 13-15 GPS satellites in view at once. Other assumptions are given in Table 11. As shown in Fig. 17, the limiting error in all components now falls below 2 cm. This opens up new possibilities for near-earth ocean altimetry, and for short-duration testing of precise instruments on the Shuttle. We should note, however, that covariance analysis can be optimistic, particularly for kinematic estimation, and unmodeled systematic errors could at least double the actual error in these examples.

CONCLUSIONS

The evidence suggests that we are obtaining a radial orbit accuracy for TOPEX/Poseidon of better than 3 cm RMS with the GPS reduced dynamic technique. Tests of orbit quality include postfit phase residuals (~5 mm), orbit overlap comparisons (~1 cm radial RMS), comparison with GSF/CPOI's (3.3 cm radial RMS; 11.5 cm maximum difference for 6 cycles), and altimeter crossovers (10 cm^2 smaller variance than GSF/CPOI or CNES orbits). The reduced dynamic orbits are seen to reduce significantly the geographically correlated error arising from the gravimetric model. Tuning of the gravimetric model with GPS data has resulted in a similar reduction of geographically correlated error in subsequent dynamic orbit solutions with all data types. Future missions can take advantage of the low cost operational GPS system developed for the TOPEX/Poseidon experiment and should obtain radial RMS accuracies of 5 cm or better in orbits as low as a few hundred kilometers.
ACKNOWLEDGMENT

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REFERENCES


## Table 1. Estimation Scenario for Dynamic Filtering of Topex/Poseidon orbit, GIPSY-OASIS II

<table>
<thead>
<tr>
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<tr>
<td>Ground Pseudorange</td>
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<td>T/P Carrier Phase</td>
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<tr>
<td>T/P Pseudorange</td>
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(all parameters are treated as constants unless otherwise specified)

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<th>Estimated Parameters</th>
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<td></td>
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<td>(cross track &amp; along track)</td>
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<td>GPS States</td>
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<tr>
<td></td>
<td>process-noise:</td>
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<td>X and Y scaling factor</td>
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<td>Y-bias</td>
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<td>Non-fiducial Station Location</td>
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<td>Tropospheric delay</td>
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<td>Carrier Phase Biases</td>
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<td>GPS and Receiver Clocks</td>
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### Table 2. Mean Coordinate Difference, GSFC POE — Dynamic

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<tr>
<th>Cycle</th>
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<td>0.604</td>
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<tr>
<td>Cycle</td>
<td>X (cm)</td>
<td>Y (cm)</td>
<td>Z (cm)</td>
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<td>--------</td>
<td>--------</td>
<td>--------</td>
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Table 3. Mean Coordinate Difference, GSFCPOE — Reduced Dynamic
Table 4: Daily Mean Difference in Z Coordinates, Cycle 18

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<tr>
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<th>Goddard/JPL, dynamic</th>
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<tr>
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<td>93mar16</td>
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<td>average, e</td>
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<tr>
<td>Orbit</td>
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<tr>
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<tr>
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Table 6. Restrictive Editing Criteria for Crossover Evaluation

<table>
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<tr>
<th>PARAMETER</th>
<th>EDIT CRITERIA</th>
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<td>Sea state</td>
<td>Significant wave height &lt; 1 m 01 &gt; 4 m.</td>
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<tr>
<td>Ocean tides</td>
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<td>Cartwright and Ray (1990), Schwiderski (1980)</td>
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<tr>
<td>Pressure loading</td>
<td>Inverted barometer &gt; 10 cm.</td>
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<td>Wind speed</td>
<td>Wind speed &gt; 10 m/s</td>
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<tr>
<td>Sea level variability</td>
<td>Mesoscale variability &gt; 12 cm (RMS)</td>
<td>Koblinsky et al. (1991)</td>
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<tr>
<td>Height interpilin.</td>
<td>Cubic spline fit RMS &gt; 5 cm</td>
<td>Aviso (1993)</td>
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### Table 7. Altimeter Crossover Statistics for Restrictive Editing Approach

<table>
<thead>
<tr>
<th>Orbit</th>
<th>No.</th>
<th>Mean (cm)</th>
<th>RMS (cm)</th>
<th>Var (cm²)</th>
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Table 8. SLR/DORIS-UTC orbit comparisons with J2'. Reduced dynamic trajectories

<table>
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<tr>
<th>Cycle</th>
<th>JGM-2 Radial RMS</th>
<th>JGM-2 3D RMS</th>
<th>TEG-3 Radial RMS</th>
<th>TEG-3 3D RMS</th>
<th>Z Shift</th>
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<td>14.2</td>
<td>13.3</td>
<td>2.2</td>
<td></td>
</tr>
</tbody>
</table>

All units cm. * indicates GPS tracking from this cycle used in gravity solution.
### TABLE 9. Error Model for Topex/Poseidon Orbit Determination Analysis

#### System Characteristics

<table>
<thead>
<tr>
<th>Parameter</th>
<th>Specification</th>
</tr>
</thead>
<tbody>
<tr>
<td>Orbit (circular)</td>
<td>1334 km, 66° inclination</td>
</tr>
<tr>
<td>Number of Ground Sites</td>
<td>6 (including 3 fiducial sites)</td>
</tr>
<tr>
<td>Number of GPS Satellites</td>
<td>18</td>
</tr>
<tr>
<td>Flight Antenna Field of View</td>
<td>Hemispherical</td>
</tr>
<tr>
<td>Flight Receiver Tracking Capacity</td>
<td>6 Channels (1.1 &amp; 1.2)</td>
</tr>
<tr>
<td>Data Types</td>
<td>1.1, 1.2 pseudorange, 1.1, 1.2 carrier phase</td>
</tr>
<tr>
<td>Data into val</td>
<td>5 minutes</td>
</tr>
<tr>
<td>Smoothed Data Noise</td>
<td>5 cm pseudorange, 1 cm carrier phase</td>
</tr>
</tbody>
</table>

#### Adjusted Parameters & A Priori Errors

<table>
<thead>
<tr>
<th>Parameter</th>
<th>Specification</th>
</tr>
</thead>
<tbody>
<tr>
<td>Topex/Poseidon Epoch State</td>
<td>1 km; 1 m/sec, each component</td>
</tr>
<tr>
<td>GPS Satellite States</td>
<td>2 m; 0.2 mm/sec, each component</td>
</tr>
<tr>
<td>Carrier Phase Biases</td>
<td>10 km</td>
</tr>
<tr>
<td>GPS &amp; Receiver Clock Biases</td>
<td>30 μsec (modeled as white noise)</td>
</tr>
<tr>
<td>Non-Fiducial Ground Locations</td>
<td>20 cm each component</td>
</tr>
</tbody>
</table>

#### Fixed Errors (Evaluated)

<table>
<thead>
<tr>
<th>Parameter</th>
<th>Specification</th>
</tr>
</thead>
<tbody>
<tr>
<td>Fiducial Site Positions</td>
<td>5 cm each component</td>
</tr>
<tr>
<td>GM of Earth Uncertainty</td>
<td>1 partin 108</td>
</tr>
<tr>
<td>Earth Gravity Error Model</td>
<td>(0-1 00% GI:M1:0-GI:M1:2 (20x20&quot;)</td>
</tr>
<tr>
<td>Zenith Atmospheric Delay Error</td>
<td>1 cm (modeled as random walk)</td>
</tr>
<tr>
<td>Atmospheric Drag Error Prior</td>
<td>10% of Total</td>
</tr>
<tr>
<td>Solar Radiation Pressure Error</td>
<td>10% of Total</td>
</tr>
<tr>
<td>Kinematic Orbit Determination Analysis</td>
<td></td>
</tr>
<tr>
<td>----------------------------------------</td>
<td></td>
</tr>
<tr>
<td><strong>Orbit (circular):</strong></td>
<td>705 km, 98° inclination</td>
</tr>
<tr>
<td><strong>Number of GPS Satellites:</strong></td>
<td>24</td>
</tr>
<tr>
<td><strong>Flight Receiver Tracking, Capacity:</strong></td>
<td>All in View (within hemisphere)</td>
</tr>
<tr>
<td><strong>Zenith Atmospheric Delay Error:</strong></td>
<td>Adjusted as Random Walk</td>
</tr>
<tr>
<td><strong>Fiducial Location Error:</strong></td>
<td>3 cm each component</td>
</tr>
<tr>
<td><strong>Earth Gravity Error Model:</strong></td>
<td>100% GEM10-GEM12 (20x20)&quot;</td>
</tr>
</tbody>
</table>
| Parameter                                      | Value
|-----------------------------------------------|-------
| Height of View (feet)                         | 320   |
| Number of Ground Stations (including 3thel 3) | 11    |
| Number of GPS Satellites                     | 24    |
| Orbit (Circular)                              | 300 km, 28° Inclination |

**Known Earth Determination Analysis**

Changes from Table 9 for Satellite
FIGURE CAPTIONS

Fig. 1 Reduced Dynamic Tracking

Fig. 2 GPS tracking system for TOPEX/POD.

Fig. 3 GPS Constellation with TOPEX.

Fig. 4 TOPEX/POSEIDON Satellite

Fig. 5 GPS Global Tracking Network

Fig. 6 Overlapping data arcs and orbit solutions

Fig. 7 Comparison of overlapping TOPEX/POSEIDON reduced dynamic orbit solutions

Fig. 8 TOPEX/POSEIDON radial reduced dynamic orbit overlaps for twelve complete 10-day cycles

Fig. 9 Comparison of TOPEX/POSEIDON dynamic orbit solutions with GPS against Goddard Space Flight Center S1 R/DORIS orbits

Fig. 10 Comparison of TOPEX/POSEIDON reduced dynamic orbit solutions with GPS against Goddard Space Flight Center S1 R/DORIS orbits
Fig. 11 Global distribution of altimeter crossovers used to evaluate orbit accuracy. A stringent editing strategy (table 6) was applied to crossovers formed from altimeter observations between January 30 and May 19, 1993.

Fig. 12. Radial amplitude spectrum dynamic minus reduced dynamic

Fig. 13 Body Fixed Z antenna offset, daily solution

Fig. 14 Body Fixed Z antenna offset, daily solution during a rare period of time in which no GPS arcs in eclipse

Fig. 15. Prelaunch Covariance Studies for TOPEX/Poseidon with 2 and 6-hr data arc lengths for a range of gravity errors. Actual RMS differences between Dynamic and Reduced Dynamic Solutions for 30-hr arcs arc shown with point a. Point x marks an estimate of the radial error in 30-hr reduced dynamic solutions.

Fig. 16. Covariance analysis prediction for future a 700 km altitude mission.

Fig. 17 Predicted error for the Space Shuttle viewing all possible GPS within a sphere

Plate 1. (a) Geographic representation of orbit height differences for GI’S-based dynamic and reduced-dynamic orbits using the JGM - 1 gravity model. A 10 X 10 spherical harmonic fit to the data captures a signal with rms amplitude of 2.3 cm.

Plate 1. (b) Geographic representation of orbit height differences for JGM-1 and JGM-2 (l’S-based dynamic orbits.)
Plate 1. (c) Geographic representation of orbit height differences for GPS-based dynamic and reduced-dynamic orbits using the tuned JGM-2 gravity model. A 10 X 10 spherical harmonic fit to the data captures a signal with rms amplitude of 1.2 cm.
Fig. 2. GPS tracking system for TOPEXPOD.
GPS Antenna and Choke Ring Back Plane

TDRSS Antenna

Radar Altimeter

Fig. 4 TOPEX/Poseidon Satellite
Fig. S. GPS Global Tracking Network
Fig. 6. Overlapping data arcs and orbit solutions
Fig. 7. Comparison of overlapping TOPEX/POSEIDON reduced dynamic orbit solutions.
Fig. 8. TOPEX/POSEIDON radial reduced dynamic orbit overlaps for twelve complete 10-day cycles.
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JGM-1 Dynamic - Reduced Dynamic

Plate 1. (a)
JGM-2 Dynamic - Reduced Dynamic

Plate 1.(c)