TOWARDS A COMPREHENSIVE MODE. OF CORONAE FORMATION ON VENUS; Suzanne E. Smrekar and Ellen R. Stofan, California Institute of Technology, Jet Propulsion Laboratory, MS 183-501, Pasadena, CA, 91109; ssmrekar@cythera.jpl.nasa.gov.

Coronae are roughly circular volcano-tectonic features that are interpreted as a manifestation of small-scale upwelling [e.g. 1, 2] and are unique to Venus. The topographic expression of coronae is highly variable, ranging from domes to plateaus, with or without moats or single or multiple outer rises [3]. Two outstanding questions in the study of coronae are how the full range of Topographic profiles are produced and the relationship between topography and the annulus of fractures that characterize coronae. Domes, plateaus, and outer rises can be formed by thermal relaxation of a topographic high due to a rising and cooling of a hot upwelling [4-7], but interior depressions, isolated rims, and high highs with rims, troughs and outer rises, can not. Relaxation can produce fractures annuli, but observed annuli frequently do not occur on the outer rise, as predicted by relaxation models [5,7]. A new model of upwelling is presented that can produce nearly the full range of observed topographic morphologies and can commonly observed off-set between tectonic fracture annuli and the outer topographic rise. The cold lithosphere at the edge of the plume head is sucked downward until the thermal anomaly dissipates, explaining the limited subduction qualities of some coronae [8,9]. This model differs from past approaches to corona formation in the use of temperature-dependent rheology and the prediction of pressure-release melting. Other aspects that may be key to the development of certain corona topographic features are the presence of a low-density depleted mantle layer beneath the high viscosity thermal lithosphere and the cooling of the thermal lithosphere during upwelling. This approach could provide constraints on thermal history.

**Model Description.** An axisymmetric finite differences scheme is used to model temperature and chemistry variations and a penalty function finite element formulation is used to solve the buoyant viscous flow equations [10]. The finite element grid is 90 by 90 elements; the finite difference grid has twice as many elements. The element spacing is non-uniform to give maximum resolution in the axial upwelling region and in the region where the plume interacts with the lithosphere. The surface temperature is 500°C. The base of the thermal lithosphere is defined as the 1100°C contour. In some models, a low density layer of mantle residuum is included beneath the thermal lithosphere. The mantle temperature is 1300°C. Viscosity is based on a dry olivine flow law and is scaled to 10^{21} Pa·s at mantle temperature. The viscosity is allowed to vary by a factor of 100 both at higher and lower temperatures, so that the viscosity varies within the temperature range of approximately 1150°C-1450°C. The dimensions of the grid are 800 km in radius and 400 km in height. The upwelling is allowed to arise from a heated region at the base of the computational domain. A radial, gaussian temperature distribution with a peak temperature of 1750°C is used to initiate the upwelling. This upwelling cools such that the peak temperature when the plume encounters the lithosphere is less than 1400°C. A similar modeling approach for venusian hotspots is described more fully in Smrekar and Parmentier [11].

**Model Results.** The evolution of the topography is shown for one upwelling model (Figure 1). In this model, the plume is allowed to form and rise for 50 my. After this point the hot region is turned off, but the plume tail continues to rise for -25 my. The thermal lithosphere is initially 50 km thick. A chemical variation is also included, in which the depletion decreases from 20% at the surface to 0% at 50 km (i.e. it is contained within the thermal lithosphere). Over the time span of the calculation, 172 my, the thermal lithosphere thickens to 110 km. In this model, the plume begins to interact with the lithosphere at about 56 my. (curve a in Figure 1). The plume tail is no longer feeding the plume by 79 my. (curve b). The lithosphere begins to thicken at the edge of the plume head as the plume spreads laterally, driving flow outward and downward. In curve c, the plume is still spreading, but the thermal anomaly is less than 50°C. The cold, dense sinking lithosphere sucks the topographic surface downward. By 116 my, after the start of the calculation, the thermal anomaly in the plume has dissipated (curve d). Once the thermal anomaly is gone, the thinned lithosphere at the center begins to sink. The viscous flow pattern and temperature distribution for this time step is illustrated in Figure 2. At 135 my. (curve e) the thermal lithosphere at the outer edge of the region is 100 km thick. As the cold lithospheric blob continues to sink into the mantle, the viscous flow sucks the delaminating region in towards the center (curves f and g), leaving a topographic depression in the center.

**Discussion.** This model illustrates a new mechanism for trench formation at coronae. Rather than being a purely viscous relaxation process, the topography is sucked downward above a sinking region of cold lithosphere. This sinking is driven initially by the viscous flow at the edge of the plume head and is sustained by the density difference between the lithosphere and mantle and the...
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difference in the thickness of the lithosphere above the plume and in the cooling lithosphere. When
the cold lithosphere thermally equilibrates, the topography rebounds. This limited sinking of the
lithosphere explains the subduction-like qualities of some coronae. In other models in which there
is a larger layer of low density depleted mantle, the low density layer is pulled downward with the
cold lithosphere. When the thermal anomaly dissipates, the low density regions rebounds
isostatically, forming a ring of high topography that is offset from the moat. These variations in the
location of the topography illustrate how the fracture annuli can be offset from the final
topography. The combination of a thick but cooling lithosphere may be essential to the formation of
at least some coronae topographic morphologies, as well as to the initiation of delamination. The
presence of a depleted chemical layer may be required to produce an isolated topographic ring, such
as seen as Heng-o Corona. This modeling approach may be able to provide constraints on the
thermal and chemical structure of the lithosphere by determining the requirements for certain
coronae planforms. Since lithospheric cooling may be required to form features such as trenches,
this represents a potentially powerful tool for interpreting tectonic history. Future work will include
matching volcanic volume estimates, scaling the models to smaller coronae, and examining the
effects of plume and lithospheric parameters.

Figure 1. Predicted topography at different time steps. See text for discussion.

Figure 2. Computational grid that corresponds to the topographic curve d in Figure 1. Shading
indicates temperature, with colder regions in darker shades. Some temperature (°C) contours are
also shown in with black lines. White lines are for 3% and 10% depletion. Arrows show viscous flow.

References: 1) S. W. Squyres et al. (1992) JGR, 97, 13,611.; 2) E. R. Stofan et al. (1992) JGR, 97,
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